



# Article Bulk Chemical and Optical Spectroscopy Characterisations of Dissolved Organic Matter Extracted from the Tropical Coastal Sediment

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Abstract: The study of organic matter characterisation in coastal sediment is useful for understanding how dissolved organic matter (DOM) reactivity influences the fate of pollutants in the aquatic environment. However, there is little information available on the structural properties of DOM in coastal sediment. Chemical analysis, UV-Visible (UV-Vis) absorbance, fluorescence, and Fouriertransform infrared (FTIR) spectroscopy were used in this work to characterise and compare the components of water-extractable organic matter (WEOM) fractions from sediment from an estuary, a coastal area, and a port in Peninsular Malaysia, Malaysia. The elemental analysis (H/C and C/N ratios) of the three samples differed, which coincided with the findings of optical indices estimated from UV-Visible and fluorescent spectroscopy analyses. WEOM had an average output of 7.05 to 8.47 mg/L and was very dominating with heavy fraction organic carbon (HFOC). In DOM-KS and DOM-K sediments, the allochthonous component with a high degree of condensation and the aromatic compound was the dominating composition. Meanwhile, the DOM-KT possessed a high autochthonous composition as well as carboxylic and phenolic content. Correlation analysis indicates that the aromaticity index SUVA<sub>254</sub> and humification index (HIX) have positive correlations. The combined results of the chemical and spectroscopic analyses indicate that different coastal ecosystems, with the integration of various human and land activities, produce variations in DOM in the coastal area.

**Keywords:** coastal sediment; water-extractable organic matter; dissolved organic matter; allochthonous/autochthonous DOM

# 1. Introduction

Coastal areas, which serve as the dynamic interface between land and marine ecosystems, play critical roles in global carbon cycling. It was estimated that 80% of the total organic carbon preserved in marine sediments is found within the deltaic and shelf regions. Coastal surface sediments serve as a major carbon burial interface, receiving huge amounts of organic carbon generated in upper waters, degrading (mainly by microorganisms), and eventually burying it [1]. Sediments, on the other hand, collect natural and anthropogenic products from the overlying water, functioning as both a sink and a source of contaminants. Coastal sediments may function as a non-point source of contaminants, especially damaging to species and human health via trophic transfer. Coastal sediment has a lower organic carbon concentration and a greater salt content than soil, making organic matter characterisation more challenging in sediment [2].

DOM has piqued the interest of academics in recent years due to its critical functions in the complexation, distribution, bioavailability, mobilisation, and immobilisation of contam-



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**Copyright:** © 2021 by the authors. Licensee MDPI, Basel, Switzerland. This article is an open access article distributed under the terms and conditions of the Creative Commons Attribution (CC BY) license (https:// creativecommons.org/licenses/by/ 4.0/). inants in the marine ecosystem [3,4]. DOM represents the active organic matter reservoir in soil, sediment and aquatic environments [5] that are composed of a mixture of different macromolecules with diverse chemical compositions, functional groups and structures [6] resulting from the degradation of natural and anthropogenic input [6]. Evidence has shown that the chemical composition of sediment DOM may vary with climate, proximity to the shore and the relative contributions of marine and continental organic matter residues [7].

Previous studies were carried out across the globe to characterise DOM, in soil, lakes, rivers and seawater [5,8–10]. Nonetheless, the data on DOM extraction from the coastal sediment, especially in the tropical region are scarce [11]. The nature and structure of aquatic organic matter are still poorly understood due to its ambiguity and heterogeneity. DOM is one of the largest pools of reduced carbon in the marine environment that is poorly characterised [12,13]. The characteristic of DOM is frequently analysed using a molecular approach via an elemental analyser and spectroscopic techniques such as UV-Vis and fluorescence spectroscopy [3,8,14].

Tropical coastal ecosystems have long been recognised for their high productivity, diverse biodiversity, and range of ecosystem services. These ecosystems, however, have undergone significant deterioration as a result of human and natural effects such as pollution, eutrophication, sedimentation, overexploitation, habitat destruction, and climate change. Peninsular Malaysia, located in the tropical area, is bordered by major waterways, including the Malacca Strait and the South China Sea. However, there has been no study pertaining to the systematic DOM evaluation for coastal sediments of Peninsular Malaysia. Because of their effects on a range of environmental processes, knowledge on the compositions of DOM is essential for understanding their reactivity with contaminants and biogeochemical processes.

The focus of this study was to describe and compare the chemical and spectroscopic features of DOM extracted from tropical coastal sediment. The findings have the potential to serve as the baseline for the establishment of the links between the physical and chemical properties and bioavailability of DOM and its possible transformation reaction.

#### 2. Materials and Methods

#### 2.1. Study Area, Sample Collection and Sample Pre-Treatment

The study was conducted in the coastal water of Peninsular Malaysia. Peninsular Malaysia is surrounded by the Straits of Malacca in the west and the South China Sea in the east, with a coastline size of 2068 km. The west coast of Peninsular Malaysia faces the Strait of Malacca which is the shortest shipping route for transportation of oil tankers from the Middle East to Asian countries such as Japan and China. As a result of heavy oil tanker traffic, it is highly susceptible to marine-based oil pollution. At the same time, increasing urbanisation and industrialisation on the west coast has resulted in high loads of land-based anthropogenic pollutants carried to the coastal areas by the rivers. The east coast of Peninsular Malaysia is situated along the coast of the South China Sea and are experiencing rapid industrial development due to petrochemical industries, agriculture and aquaculture activities and urban development. In addition, major ports (i.e., Kuantan Port) a multi-cargo port has served as an important seaway to major ports in Asia-Pacific regions [14].

In this study, three surface coastal sediment samples were collected from three coastal environments namely: Kuala Selangor (KS), Kuala Terengganu (KT) and Kuantan (K) (Figure 1). Sampling activity was conducted between March 2016 and June 2016. The sediment samples were collected by using Smith McIntyre grab sampler, transferred into plastic containers, and stored at 4 °C and then transported to the laboratory. For every sampling station, three sub-samples were collected. Details of the sampling stations were described in Table 1.



Figure 1. Location of sampling stations in this study.

<b>Fable 1.</b> Description of sampling static
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Station	Longitude	Latitude	Depth of Water Column (m)	Description
Kuala Selangor (KS)	101°18′05.5″ E	3°13′48.5″ N	15 m	3 km from estuary with mangrove forest Fishing village, aquaculture pond
Kuala Terengganu (KT)	103°09′23.2″ E	5°20′36.9″ N	15 m	5 km off Kuala Terengganu river mouth
Kuantan (K)	103°22'39.2″ E	3°48′28.3″ N	15 m	2 km of Kuantan port

# 2.2. Bulk Marine Sediment Analysis

In the laboratory, the wet sediment samples were air-dried at room temperature and later sieved through 100 mesh screens to eliminate coarse impurities such as rocks, plant matter, and mollusk shells. The marine sediments were air-dried in a 60 °C in an oven until they reached a constant weight and then homogenized in an agate mortar. Sediments pH was measured at a 1:2.5 (w/v) ratio by using Hanna 211 pH meter. The sediment texture and specific surface area were analyzed using a Laser Diffraction Particle Size Analyzer (SALD-2300, Shimadzu, Japan)

The carbonate content was removed according to the method described by [15]. Briefly, 3.0 g of each sample were air-dried and mixed with 10 mL of 1 M HCl and left for 24 h. Then, the samples were centrifuged for 1 min (3000 rpm) and the supernatant discarded. The procedure was repeated until no further effervescence in the sample. The samples were rinsed and centrifuged three times with ultrapure water, dried below 50 °C, and weighed. Total organic content (TOC) analysis was performed by modified Walkley–Black wet combustion method as described by [16]. This method consists of titration of dichromate with ferrous ammonium sulfate, which is obtained after wet oxidation of the samples with potassium dichromate using ferroin as an indicator.

Light fraction organic carbon (LFOC) was extracted according to the method describes by [17]. Briefly, 5.0 g of the samples were shaken in NaI for 16 h, and the suspensions were allowed to settle for 72 h. Subsequently, the supernatant in the top 1 cm of the NaI solution was transferred by suction to a Millipore filtration unit and was thoroughly washed with distilled water to remove NaI. After oven drying at 100 °C, LFOC was weighed, and the heavy fraction organic carbon (HFOC) was calculated by the difference between TOC and LFOC. Determination of active organic carbon (AOC) was based on the method as described by [18]. Generally, KMnO<sub>4</sub> solution was added to 0.5 g of dried sediment and undergoes agitation at 200 rpm for 2 h, and followed by centrifugation at  $2000 \times g$  for 5 min. Then, the solution was diluted and measured spectrophotometrically at 565 nm. The concentration of AOC was calculated according to the difference in the KMnO<sub>4</sub> concentration before and after oxidation. Meanwhile, total acidic group, carboxyl and phenolic content were determined according to the procedures described by [19]. Total acidic and carboxylic group contents were determined by the  $Ba(OH)_2$  titration and  $Ca(AC)_2$  titration methods, respectively. The total phenolic group amounts were then calculated by subtracting the carboxyl group contents from the total acidic group contents.

#### 2.3. Extraction of DOM from Marine Sediment

DOM was extracted from sediment samples by the water extraction method according to the procedure recommended by [20] with slight modification. Briefly, a total of 10.0 g of the air-dried sample was added in 100 mL of ultrapure water for and the sediment suspension was agitated for 24 h, 28 °C, at 150 rpm and purged with N<sub>2</sub> gas to minimise the oxidation process. The sediment-water mixture was then centrifuged for 20 min at 5000 rpm, and the supernatants were filtered through 0.45 µm pre-combusted GF/F filters and stored at 4 °C in an amber vial. DOM concentration was quantified by using a total organic carbon analyser (Shimadzu, L-<sub>CPN</sub>). The pH of the DOM was adjusted to ≈3 prior to UV-Vis and fluorescence spectroscopy analysis. DOM samples were stored at 4 °C until further analyses. The elemental analysis of the DOMs was determined in triplicate by using a Vario EL 3 Elemental Analyzer (Elementary, Germany).

#### 2.4. UV-Vis Spectroscopy Characterization

Prior to UV-Vis spectroscopy analysis, the dissolved organic carbon (DOC) concentration of all DOM samples was diluted to 10 mg/L with ultrapure water to minimize the inner filtering effects [21]. The UV-Vis absorbance of each extracted DOM was measured by using a UV-Vis spectrometer (Orion Aquamate 8000 UV-Vis spectrophotometer, Thermo Scientific) within a spectrum of 250–600 nm, at 1 nm increments, using a 1 cm quartz cuvette. Several spectral indices were used for this study including the SUVA<sub>254</sub>, spectral slope ratio (S<sub>R</sub>) the relative slope over 275–295 nm to 350–400 nm), the ratio of the absorptions at 250 and 365 nm ( $E_2/E_3$ ), the ratio of absorptions at 465 and 665 nm ( $E_4/E_6$ ) and the index for CDOM abundance (A<sub>355</sub>).

## 2.5. Fluorescent Spectroscopy

Synchronous fluorescence spectra were measured with a luminescence spectrometer (Perkin Elmer  $LS_{55}$ ). The wavelength range of excitation and emission were set from 220 to 550 nm and from 280 to 550 nm, respectively, at 0.5 nm increments. Humification

index (HIX), fluorescent index (FI), and biological index (BIX) were included for the fluorescence indicators [22,23]. HIX is a ratio of the areas under the emission spectra over 435–480 nm to 300–345 nm with an excitation wavelength of 255 nm. FI is a ratio of the fluorescence intensities at the emission wavelength DOM of 450 and 500 nm with an excitation wavelength of 370 nm [4]. BIX is a ratio of the fluorescence intensity at the emission wavelength DOM of 380 nm and 430 nm with the excitation wavelength of 310 nm [24].

#### 2.6. Attenuated Total Reflectance Infrared ATR-FTIR Analysis

DOM samples were freeze-dried to obtain powder samples for further ATR-FTIR analysis. ATR-FTIR spectra were obtained using an FTIR spectrometer (Shimadzu, IRTracer, 100). Five hundred and twelve scans were co-added to obtain each spectrum with a resolution of  $4 \text{ cm}^{-1}$  in the 400–4000 cm<sup>-1</sup> wavenumber range.

#### 2.7. Statistical Analysis

The data are presented as mean with standard deviation. Statistical test was carried out using SPSS Version 18.0 (SPSS Inc., Chicago, IL, USA). Descriptive statistical analysis was carried out to describe the main characteristic of marine sediment collected from the three locations. In addition, analysis of variance (ANOVA) was used to compare the means with respect to the three categorized sampling locations. The spectral slopes were fit by using Sigmaplot Version 14 statistical software.

#### 3. Results

## 3.1. Properties of Bulk Marine Sediment

In this study, three surface marine sediment samples were characterised, and the properties of the sediment are shown in Table 2. The pH values did not vary greatly (between pH 7.45 and 7.81) among the samples and the lowest pH value (KS) may be attributed to the oxidation of iron sulfides. The pH values are in the range of values as reported by [14].

Station	TOC(g/kg)	<b>Grain Distribution %</b>		tion %	Sediment	Total Acidic Group	COOH Group	Phenolic Group	nН
Station	Mean ( $\pm$ SD)	Clay	Silt	Sand	Classification USDA	(mmol/g)	(mmol/g)	(mmol/g)	PII
KS	$1.25\pm0.73$	46	20	34	Clay	7.34	4.12	3.22	7.45
KT	$0.83\pm0.30$	10	24	66	Sandy Íoam	6.52	3.60	2.55	7.56
K	$1.08 \pm 1.04$	37	31	32	Clay loam	6.23	3.50	2.73	7.81

## Table 2. Properties of the marine sediment.

TOC is an important index to study organic matter content and has been widely applied to modern sediments [25]. In this study, TOC was varied in between 0.83-1.25% with an average value of  $1.05 \pm 0.1\%$  which is in good accordance with the concentration ranges reported in other coastal areas in Malaysia [26]. A higher level of TOC was observed in KS and K, and these sediments were dominated by fine-grained sediments. On the other hand, sand sediment was found in KT, give the lowest value of TOC. TOC measurements revealed positive linear correlations between fine-grained percentage and sedimentary TOC [27] due to the large specific surface area of fine-grained sediment [28], which provides more binding sites for the adsorption of organic matter. Furthermore, the sandy sediment has greater penetration of oxygen, compared to fine-grained and muddy sediment; this may help to explain the relatively lower value of TOC KT compared to KS and K. Furthermore, the KS received water input from Selangor River, therefore considering the backwater effects of the Selangor River, organic matter was easily deposited, resulting in high organic carbon content in the sediments.

The sediment from KS was highest in the total acidity group, with the value of 7.34 mmol/g, whereas the values for KT and K were 6.52 and 6.23 mmol/g respectively. The aforementioned differences in acidity properties may affect their ability to bind pollutants

as the acidic functional group content is an important factor controlling complexation by organic matter ligands [29]. According to [30] DOM derived from the aquatic system contained a higher acidic functional group than the soil.

## 3.2. Elemental Analysis of DOM

Element analysis is one of the most common methods to record DOM structures. Evidence has shown that atomic ratios of carbon/nitrogen (C/N), hydrogen/carbon (H/C), oxygen/carbon (O/C), and are frequently used to identify DOM from different sources and to elucidate structural formulas for DOM [31]. Ranges of the elemental composition of the isolated DOMs as well as the calculated ratios of the H/C, O/C and C/N, are shown in Table 3.

Table 3. Elemental composition and atomic of DOM extracted from three marine sediments.

Station	C (%)	H (%)	N (%)	O (%)	H/C <sup>a</sup>	O/C <sup>b</sup>	C/N <sup>c</sup>
KS	28.60	30.20	1.60	39.61	1.06	1.38	17.88
KT	30.40	15.01	3.52	48.64	0.49	1.58	8.69
Κ	33.10	35.83	2.72	27.53	1.08	0.90	11.04

<sup>a, b, c</sup> Calculated as described by [8].

As can be seen in Table 3, variation of elemental composition was assessed in the DOM samples. The ratio values of H/C increased in the order of DOM-KT < DOM-KS < DOM-K, which correspond to 0.49, 1.06 and 1.08 (mean average  $0.88 \pm 0.04$ ), respectively, which indicates the increasing ratio of aliphatic carbon [32]. The H/C values were found greater than 1 in KS-DOM and K-DOM, indicates these samples probably originate from vascular plant material instead of algal or bacterial [8]. Differently, the H/C value of DOM-KT was 0.49, this value is characteristic of organic matter with autochthonous origin.

Further confirmation was made by calculating the C/N ratio. Allochthonous vascular plant materials contain cellulose, lignin, and tannin [6] and have C/N ratios of more than 12, while the non-vascular C/N ratio is between 2 to 10 [7] and those intermediate values (10–12) suggest a mixture of allochthonous and marine organic matter. Thus, considering that the C/N ratio values obtained for DOM-KS and DOM-K ranged from 15.10–18.40 (mean 17.88) and 10.20 to 11.4 (mean 11.04) and the values found for the DOM-KT varied from 7.66 to 9.20 (mean 8.69), we can suggest that DOM-KS is allochthonous and DOM-K is a mixture of allochthonous and autochthonous, and DOM-KT is autochthonous. In a study conducted by [33] in the Batan Bay Estuary, Philippines, it was stated that the high C/N values also imply a persistent input of fresh organic matter from the residual mangrove stands and low SOM decomposition and organic matter preservation. The O/C ratio was relatively high for DOM-KT as compared to those DOM-KS and DOM-K. This indicates a high composition of carboxylic and O-alkyl groups [31] and implies a greater presence of oxygen-containing functional groups, [5] in DOM-KT.

# 3.3. DOC in Extracted DOM

To quantify the DOM in marine sediment, DOC concentration is generally used to represent the content of DOM [34]. The value of DOC and variations of organic carbon (OC) content in the studied sediment can be found in Supplementary Materials (Table S1) and Figure 2a,b. DOC values were found to be in the range of 7.05–8.47 mg/L. By the one-way ANOVA, the difference in DOC concentration was significant between the three samples (p < 0.05). The highest DOC concentration was observed in DOM-KS and the lowest DOC concentration was observed in DOM-KT.



**Figure 2.** Variation of OC in bulk marine sediment, (n = 9). (**a**) Concentration of DOC and CDOM, and the CDOM/DOC ratio. (**b**) Percentage of LFOC, HFOC and AOC in bulk marine sediment. Error bars indicate the standard deviation.

Further characterisation of the organic carbon was determined through the LFOC, HFOC and AOC content (Figure 2b). Generally, the LFOC ranged from 0.76 to 1.30 g/kg and HFOC value from 7.56 to 11.2 g/kg. LFOC material consists of incomplete decomposed organic residues and serves as an indicator of the effects of cropping practices [35]. HFOC contents were 89.6 to 91.7% of the TOC, indicating the dominant fraction in all three samples. These findings were in agreement with a previous study by [18]. The AOC represents the component of the organic carbon that can be readily decomposed by microbes and can supply direct nutrients ranged from 1.43 to 2.21 g/kg. The LFOC, HFOC and AOC were significantly higher in KS than in K and KT (p < 0.05).

#### 3.4. UV-Visspectroscopy Analysis

The DOM's UV-Vis absorption was attributed to its unsaturated bond and aromatic composition. The absorption or ratio of UV-Vis at specific wavelengths has therefore been widely used to investigate the structural characteristics of DOM, such as the content of aromatic compounds, the degree of polymerisation and the molecular weight [36]. The DOM UV-Vis absorption spectrum (Figure 3) shows an exponential decrease with increasing wavelengths, following a pattern similar to that already described for the sample of natural organic matter [37]. In addition, the spectrums exhibited a shoulder at around 250 nm to 280 nm and were mainly attributed to aniline derivatives, phenolics, aromatic carboxylic, and polycyclic aromatic compounds ( $\pi$ - $\pi$ \* transition) [38].

The additional analyses of SUVA<sub>254</sub>,  $E_2/E_3$  ratio,  $E_4/E_6$  ratio,  $A_{355}$  and  $S_R$  indices were used to further study the chemical characteristics of DOM extracted from different marine sediment. The calculated UV-Visiindices of the DOM were represented in Table 4. The average SUVA<sub>254</sub> values in DOM-KS was significantly higher compared to other two DOM sample. The SUVA<sub>254</sub> ranged from 4.19 to 5.44 (mean value 4.79  $\pm$  0.53) indicates that the DOM sources are derived from vascular plants enriched with aromatic components [39]. The SUVA<sub>254</sub> are known as aromatic and molecular weight tracers and are positively correlated with the contents of aromatic carbon [40] and the presence of unsaturated double bonds in the conjugate [41]. The results are consistent with the findings of others [42,43] related to hydrophobicity and hydrophilicity of the DOM sample.



Figure 3. UV-Visible absorbance spectra of DOM of KT, K and KS in this study.

Table 4. UV-visible and fluorescent spectroscopy indices of the DOM extracted from three sediments.

Station -		U	V-Visible Indic	Fluorescent Indices				
	SUVA <sub>254</sub>	$E_2/E_3$	E4/E <sub>6</sub>	A <sub>355</sub>	S <sub>R</sub>	HIX	FI	BIX
KS	$5.44\pm0.08$	8.32	4.32	3.43	0.13	$1.63\pm0.03$	$0.95\pm0.08$	$0.67\pm0.02$
Κ	$4.75\pm0.09$	10.71	6.50	2.77	0.07	$1.36\pm0.02$	$1.10\pm0.07$	$0.62\pm0.01$
KT	$4.19\pm0.08$	18.43	8.64	3.98	0.59	$1.20\pm0.02$	$1.58\pm0.09$	$1.27\pm0.12$

The values shown as mean with (n = 9), followed by  $\pm$ (standard deviation).

Hydrophilic fractions were associated with low SUVA<sub>254</sub> (<3) and hydrophobic fractions of supposedly higher aromaticity were associated with low SUVA<sub>254</sub> (>4) [43]. This suggests that samples from DOM-KS had the highest DOM aromaticity and hydrophobic fraction in comparison with other samples. This could be caused by the decay of allochthonous organic substances material that contains various classes of aromatic compounds [6].

The mean molecular size of DOM derived from the  $E_2/E_3$  ratio were 8.32, 10.71 and 18.43 for DOM-KS, DOM-K and DOM-KT, respectively. The aforementioned decreasing ratios are because of increasing molecular size [44]. As the molecular size increased,  $E_2/E_3$  decreased because of stronger light absorption by high-molecular-weight (HMW) DOM at longer wavelength DOM [45]. Meanwhile, the  $E_4/E_6$  ratio, the indicator of the degree of condensation, was highest in KT, consistent with the earlier findings of H/C ratio and SUVA<sub>254</sub>, confirming the presence of a higher aliphatic compound in DOM-KT than in DOM-KS and DOM-K.

The adsorption coefficient at 355 nm (A<sub>355</sub>) was selected as the index for measuring chromophoric organic matter (CDOM) abundance [24,46] due to the complexity of DOM. The highest value of A<sub>355</sub> was observed in DOM-K, indicates that the highest abundance of CDOM. In addition, the CDOM/DOC ratio was calculated to estimate the relative contribution of CDOM to the bulk DOC (Figure 2a). The average CDOM/DOC (L mg-C<sup>-1</sup> m<sup>-1</sup>) was 0.42 (DOM-KS), 0.35 (DOM-K) and 0.55 (DOM-KT) (Figure 2a), respectively, showing the relative contribution of CDOM was greater in KT. A strong linear relationship between CDOM and DOC (Figure 4) in all DOM (R<sup>2</sup> = 0.99) suggests that more than 90% of DOC variability among the DOM can be explained by variations of the quantity of



CDOM [46]. The strong linear relationship of CDOM and DOC was consistent with a previous study [47,48].

**Figure 4.** Relationship between DOC and CDOM in KS, K and KT, *n* = 9.

According to [40],  $S_R$  of  $S_{275-295-350-400}$  nm in the ultraviolet absorption spectrum was an indicator to measure the molecular weight and source of DOM. The larger the  $S_R$ , the smaller the molecular weight. It can be seen from Table 4 that the  $S_R$  values of all DOM samples were less than 1.0, ranged from 0.13 to 0.59 with a mean value of  $0.54 \pm 0.51$  indicating again that high molecular weight molecules are dominant [23] which is consistent with the result obtained from SUVA<sub>245</sub> and  $E_2/E_3$  ratio.

# 3.5. Fluorescent Properties of DOM

Fluorescence spectroscopy provides a specific evaluation of DOM in which only fluorophores are analysed, unlike UV-Vis spectroscopy that detects any compound that absorbs light (chromophore) [3]. The HIX is commonly used as an indicator of the DOM age and recalcitrance [5]. The higher the HIX value, the more difficult the DOM can be degraded and used by the microorganisms, and the lower the molecular activity is [49].

Similar to SUVA<sub>254</sub>, DOM-KS showed a significant highest HIX value (p < 0.01), which implied that the presence of more complex organic and aromatic material than DOM-K and DOM-KT. As depicted in Figure 5 and Table 4, the HIX values of all the three samples ranged from 1.21 to 1.63 (mean value  $1.39 \pm 0.18$ ) and which are in the range of the previously published study [5,14,50]. However, HIX values were less than 4.0 and this result indicated that the humification degrees of DOM in DOM-KS, DOM-K and DOM-KT were weak. Our data show that there is a strong significant correlation between HIX and SUVA<sub>254</sub> (Figure 6) as described by other studies in coastal areas.



Figure 5. Fluorescent spectral indices of DOM in this study. (a) HIX (b) FI (c) BIX.

Notably, FI can be used to differentiate between allochthonous and autochthonous DOM sources as well as data on the source of DOM. FI is an indicator of DOM source with higher (~1.4) and lower (~1.2) values indicating the microbial and terrestrial sources, respectively [4]. It was found that FI values for all three samples were in a range of 0.95 to 1.58, in which the lowest value was in DOM-KS, corresponding to the terrestrial input (allochthonous). By contrast, the FI value for DOM-KT is 1.58 and this indicates the microbial sources. This finding indicates that the productivity and self-purification ability of KT sediments are strong, which is inconsistent with the good water quality at present [50].

There was a significant difference between the BIX value for KS with KT and K (p < 0.05). This demonstrates the possibilities of a different magnitude of an autochthonous component among the sediment samples. Furthermore, the values of BIX obtained from this study range from 0.67 to 1.27, indicating high terrigenous component in DOM-KS and DOM-K and autochthonous component in DOM-KT.



**Figure 6.** Pearson Correlation of HIX and SUVA<sub>254</sub>, n = 27.

The results are similar to those reported for sediment estuarine [3,51]. A BIX value greater than 1, explains the increase in autochthonous components, whereas a low BIX value between 0.6 and 0.7 showed terrigenous input or human impacts are large [36]. Since there are significant differences of BIX value between the DOM (p < 0.05), it explained the possibilities of a different magnitude of an autochthonous and allochthonous component among the DOM samples. Moreover, evidence has shown that BIX is used to determine the presence of  $\beta$  fluorophore, which is characteristic of autochthonous biological activity [3].

## 3.6. DOM ATR-FTIR Spectra

Figure 7 shows the ATR-FTIR spectra of the three DOM studied in this study. In general, the spectra of all three locations displayed a similar peak pattern with ten distinctive adsorption bands but with different absorbance intensities. The major bands observed in the spectra are in agreement with results reported by other researchers [52,53].

The typical broad bands with peaks around  $3400 \text{ cm}^{-1}$  to  $3000 \text{ cm}^{-1}$  which is the broadest band were attributed to O–H stretching of alcohols and/or phenols and N–H stretching of amines and/or amides [8]. The band regions around 2920 cm<sup>-1</sup> and 2850–2856 cm<sup>-1</sup> are represented aliphatic C–H vibrations of aliphatic methyl and methylene groups [54]. The intensities of those peaks (2920 and 2850) were higher for KT and K than KS, thus illustrating that the KT and K contained more aliphatic compounds than the KS. The strong adsorption bands occurring at 1650–1630 cm<sup>-1</sup> are attributed to aromatic C=C skeletal vibrations, C=O stretching and N-H bending. A peak at 1550–1520 cm<sup>-1</sup> belongs to the aromatic C=C stretching of lignin residues and a peak at about 1460 cm<sup>-1</sup> is attributed to aliphatic C–H. A peak between 1235–1220 cm<sup>-1</sup> is assigned to lignin composition [55]. A band that appears around 1075 cm<sup>-1</sup> is attributed to C–O stretching of polysaccharides and polysaccharide-like substances, and/or Si–O of silicate impurities [56]. Finally, some bands appeared below at 700 cm<sup>-1</sup> and these bands are assigned for multiple functional groups/compounds such as sulfate, carbohydrate, alkyl-halide, halosilanes, nitro group and C–S bonds [53].



Figure 7. ATR-FTIR spectra of DOM samples.

FTIR spectroscopy provides information on the relative distributions of functional groups in the DOM. The results obtained suggest that all DOM consists of alcohol of carboxyl and compounds, and some amines and nitrates. Conversely, two sharp and strong peaks at 2920 cm<sup>-1</sup> to 2850 cm<sup>-1</sup> indicate that all DOM sample comprises a large amount of aliphatic. The intensity of the absorption bands depends on the amount of absorbing functional groups. This means that larger contents of functional groups result in greater intensity of the corresponding absorption bands, whereas smaller contents result in less intensity. It can be seen from Figure 7 that DOM-KT have highest adsorption intensity at 1650 to 1630 cm<sup>-1</sup> compared to other DOM, reflects than this sample contains more C=O compound.

# 4. Discussion

Characterisation of organic matter in coastal ecosystems enables us to understand the biogeochemical transformation of organic matter and its influence on ecosystem productivity in response to various changing environmental conditions [57]. It is generally accepted that the variation of organic matter in coastal sediment is due to the external input and internal decomposition process. This is mainly because of the strong diffusion of DOC in the oxidising environment and the decomposition of a large number of vegetation litter on the surface, which results in higher organic carbon content in the surface of sediments, and then gradually decreases with the increase in depth [58]. Different fractions of organic matter are known to regulate the behaviour of contaminants such as trace elements in the water ecosystem [59].

In this study, the KS sediment sample generally shows the highest content of acidic functional groups compared to KT and K samples (Table 2), thus suggesting that the KS had the highest chemical reactivity and more binding sites than KT and K samples. High

phenolic content in sample KS may be due to the presence of lignin phenolic sources [56]. The DOM was extracted by the water extraction method and the concentration was then recorded as DOC. Our results show that the DOC concentrations were relatively higher compared to the DOC in the water column than or similar to those recorded for the estuary water column [60]. This might be explained due to the dilution factor of a large volume of water in the water column. Furthermore, in natural conditions, the inflow of coastal water into the coastal ecosystem and the influence of coastal winds may cause sediment resuspension and later further increase DOM and nutrient concentration in the water column [61].

Studies of sediment in several coastal ecosystems revealed that the organic matter consisted of a mixture of multiple sources from terrestrial sources and marine sources such as microalgae [13,62]. The sources were usually divided into autochthonous and allochthonous. The autochthonous source was derived from the decomposition of microorganisms, plants, phytoplankton or algae residue and the extracellular enzymes secreted by microorganisms. Meanwhile, allochthonous organic matter including terrestrial materials transported via rivers and marine materials delivered through tidal actions. Some researchers used biomarkers such as lipid biomarkers [1], fatty acid and biopolymers biomarkers [63] in studies of organic matter. Our findings have shown that there are variations in the sources of marine sediment DOM from different backgrounds, depending on land activities, which could be associated with land-derived input and human activity.

DOM-KS site is located nearby Kuala Selangor estuary which receives large amounts of land-based pollutants from its developed surroundings in the form of runoffs and also via atmospheric deposition [64]. As mentioned before, the H/C and C/N ratio values are relatively high, confirming the hypothesis that vascular plants are the dominant contributors in DOM-KS. Studies in the literature showed that the high organic matter in the estuary was contributed by the decomposition of root material and decomposition of leaves of the mangroves forest [65]. The contribution of mangrove-derived organic matter to the sediment was also influenced by a variety of factors including mangrove density, hydrodynamic condition, geomorphological setting and the distance from the mangrove [66]. Therefore, we speculate that major input from the mangrove-derived organic matter contributed the highest influence for high TOC and DOC values in KS. High organic matter content in KS might also be explained by the aquaculture farming activity available in the area. A recent study in Monastir Bay, Tunisia indicates the high levels of sediment organic matter are due to excretions of fish and faecal materials from fish farming activity [67].

In a context of intensive anthropogenic action, the K sampling site which is situated nearby Kuantan port is exposed to heavy industrial, port and urban pressure. The port has served as an important seaway to major ports in Asia-Pacific regions [68]. Furthermore, ports are low-energy areas with high rates of accumulation of organic contaminants in the sediment [60]. Moreover, it was previously reported that the use of paints and grease repellents for ship and dock protection could contribute to organic matter deposition on the sediment. The accumulation of diverse sources derived from organic carbon and nitrogen (natural and anthropogenic origin) in the sediment is also influenced by the hydrodynamic condition of the system. In [63], it was emphasised that sampling sites subjected to low hydrodynamic conditions favour the deposition of organic matter-rich fine particles in contrast to sites subjected to strong hydrodynamism. This statement is applicable to the port and harbour ecosystem.

In contrast to KS and K, it is interesting to observe that TOC values of DOM-KT were the lowest among the three sampling stations. Previously [69] reported the source of organic carbon in the coastal sediment of Kuala Terengganu is due to aromatic hydrocarbon with evidence of the presence of gasoline, diesel, and mineral oil that can be linked to oil spills from fishing vessels and high-speed boats. The area is also facing the South China Sea and is vulnerable to oil pollution because of intense petroleum production activities

in the area. The O/C ratio from the elemental analysis was relatively high for DOM-KT, suggesting the organic matter is rich with carboxylic and phenol content.

In comparison with UV-Vis absorption spectra, FTIR analysis is used to reveal the presence of specific functional groups. Information on the active functional group in DOM is important to understand the pollutant's behaviour in the ecosystem. The active functional groups such as carboxyl, carbonyl, and hydroxyl, can interact with several metal ions, organic matter, and toxic reactive pollutants in the overlying water and sediments through adsorption, complexation, oxidation, and reduction. Thus, they affect the chemical forms, migration and transformation, bioavailability, and destination of toxic reactive pollutants. These interactions and the structure and function of whole aquatic ecosystems may have a key role in the decomposition or elimination of toxic or harmful substances and the balance of nutrient elements in the water environment [50]. The ATR-FTIR spectroscopy data suggest that the DOM's contained organic polyelectrolytes with high contents of carboxyl, carbonyl, and hydroxyl group (Figure 7). The differences of DOM-KT and the other two DOM can be observed specifically at O-H stretching region (3400–300 cm<sup>-1</sup>) and C=O stretching region (1650–1630  $\text{cm}^{-1}$ ). The DOM-KT spectra show the highest intensity in these two regions compared to the other DOM. Carboxylic and phenolic substituents are associated with terrestrial contributions [70]. The FI and BIX values of DOM-KT from the fluorescent spectroscopy analysis indicate a predominantly autochthonous origin freshly produced from biological or microbial origin. This is supported by the absence of lignin peak (1220–1235 cm<sup>-1</sup>) in ATR-FTIR in the DOM-KT sample. Lignin which is a vascular phenolic polymer had been used as a biomarker for terrigenous DOM in other studies [71,72].

To better understand the characteristics of each DOM, we analysed the relationship between spectroscopy indices through Pearson correlation analysis (Table 5). A significant positive relationship was observed between the SUVA<sub>254</sub> and HIX. However, there was a significant negative correlation with other spectroscopy indices excluding the A<sub>355</sub>. As described by [43], SUVA<sub>254</sub> is positively correlated with humic acid fraction and inversely correlated with average molecular weight ( $E_2/E_3$ ). There was no significant correlation between the A<sub>355</sub> and HIX and A<sub>355</sub> and E<sub>4</sub>/E<sub>6</sub>.

	SUVA <sub>254</sub>	$E_2/E_3$	$E_4/E_6$	A <sub>355</sub>	HIX	BIX	FI
SUVA <sub>254</sub>	1						
$E_2/E_3$	-0.926 *	1					
$E_4/E_6$	-0.984 *	0.946 *	1				
A <sub>355</sub>	-0.316	0.606 *	0.339	1			
HIX	0.979 *	-0.905 *	-0.982 *	-0.235	1		
BIX	-0.878 *	0.946 *	0.895 *	0.627 *	-0.825 *	1	
FI	-0.904 *	0.957 *	0.906 *	0.597 *	-0.866 *	0.898 *	1

Table 5. Correlation among spectroscopy indices.

\* Indicates significant correlation at the 0.01 level. Correlation is significant at the 0.01 level (2-tailed).

## 5. Conclusions

The findings of this study revealed that the chemical and spectroscopy characterisation of DOM isolated from the tropical coastal marine sediment was influenced by the physical environment and land activity.

The TOC of sediment samples was varied between 0.83-1.25% (w/w) and was dominated with HFOC. The DOM concentration (measured as DOC) was found to be in the range of 7.05–8.47 mg/L. Further analysis of elemental ratio shows the H/C ratio in KS-DOM and K-DOM were more than 1, which indicates the terrestrial derived sources mainly from mangroves-derived organic matter and anthropogenic activities. Meanwhile, the lower value of H/C in DOM-KT (less than 1.0) revealed it is rich in autochthonous marinederived organic matter. This was confirmed with the C/N ratio value and several UV-Vis and fluorescent spectroscopy indices. The UV-Vis spectra analysis via the SUVA<sub>254</sub>,  $E_2/E_3$  ratio and  $S_R$  revealed that high molecular weight molecules are dominant in all DOM samples. Furthermore, the ATR-FTIR spectrum confirmed the presence of functional groups such as organic polyelectrolytes with high contents of carboxyl, carbonyl, and a hydroxyl group.

Future studies should be extended to examine the molecular composition of the organic matter from the coastal sediments. A 3D-EEM fluorescence-PARAFAC characterisation would provide additional information on the DOM diversity and biochemical constituents in the marine sediment. The method provided for DOM characterisation of the KS, KT, and K might be extended to other tropical coastal ecosystems that have received less attention, such as wetland, coral reef, and peat coastal regions. This assessment will be valuable for future comparative organic matter profiles in tropical ecosystem sediments, as well as for updating the current emerging pollution profile in this anthropogenically damaged area.

**Supplementary Materials:** The following are available online at https://www.mdpi.com/article/ 10.3390/jmse9090997/s1, Table S1: Variation of organic carbon (OC) content in the three sediments.

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